BOSE-EINSTEIN APPROXIMATION IN SELF-CONSISTENT EQUATION OF STATE OF MINERALS

P.I.Dorogokupets

Institute of the Earth's Crust, Irkutsk, 664033, Russia

This study was supported by the Russian Fundamental Research Foundation (Project No. 99-05-64891) Herald DGGGMS RAS № 5 (15)'2000 v.2

URL: http://www.scgis.ru/russian/cp1251/h dgggms/5-2000/term2.eng

Introduction

Knowledge of self-consistent thermodynamic functions of minerals is the main requirement for computer modeling of chemical equilibria and construction of geophysical models in lower crust and mantle of the Earth. This problem is resolved in well known method of potentials in the Debye approximation [e.g. 1-6]. Other approaches to this problem are known also [e.g. 7-10]. Otherwise for calculation of the thermodynamic functions vs. T, P and V it is necessary to enter different simplifications as it was made in *Anderson and Zou* [11] formulation of the thermodynamic functions for mantle minerals.

The main purpose of this study is the modification of the Zharkov and Kalinin [1] formalism. In place of the usual Debye function, we are using the Bose-Einstein functions for the approximation of thermal part of the Helmholtz free energy. The Bose-Einstein functions was proposed by Kut'in and Pyadushkin [12] for the Gibbs energy at zero pressure, but these relations may be used in the VT relations, that is for the Helmholtz free energy. As a result we obtain a set of equation for volume-temperature dependence of a thermodynamic functions. We can apply the derived set of equations for calculation of any thermodynamic functions vs. PVT if we find a set of fitting parameters in our experimental model optimizing available measurements of the heat capacity and relative enthalpy vs. T, volume as function of pressure and temperature, thermal expansion coefficient, bulk moduli etc. Finally, we obtain a simple thermal equation of state for mantle minerals which one evaluate allows both to self-consistent thermodynamic functions from experimental measurements and to calculate one vs. any temperature and pressure. This EoS is equally suitable both for petrological constructions and for geophysical models of the deep of the Earth.

Basic equations

The Helmholtz free energy F(V,T) may be written as [1, 13]:

 $F(V,T)=U_0+E_P(V)+F_{th}(V,T)+F_a(V),$ (1) where U_0 is energy at 0 K and zero pressure, $E_P(V)$ is a part of energy depend only from volume, $F_{th}(V,T)$ is a thermal part of the free energy, $F_a(V)$ is anharmonic part of the free energy.

Energy vs. volume may be expressed using the logarithmic equation of state by *Poirier and Tarantola* [14]:

$$E_P = \frac{1}{2} K_0 V_0 [(\ln x)^2 - \frac{K' - 2}{3} (\ln x)^3], \tag{2}$$

where K_0 and V_0 are bulk modulus and volume at zero temperature and pressure, $K'=dK_0/dP$, $x=V/V_0$. The pressure $P(V)=-\partial E_P/\partial V$ and bulk modulus $K_T(V)=-V(\partial P(V)/\partial V)$ written as:

$$P(V) = K_0 \frac{1}{x} \left[-\ln x + \frac{K' - 2}{2} (\ln x)^2 \right], \tag{3}$$

$$K_T(V) = K_0 \frac{1}{x} [1 + (1 - K') \ln x + \frac{K' - 2}{2} (\ln x)^2],$$
 (4)

The thermal part of free energy expressed using the Bose-Einstein functions [12]:

$$F_{th} = m_B R T \left[\frac{\Theta_B}{2T} - \ln(1+b) \right] + m_{E1} R T \left[\frac{\Theta_{E1}}{2T} + \ln(1 - \frac{1}{\exp(\Theta_{E1}/T)}) \right] + m_{E2} R T \left[\frac{\Theta_{E2}}{2T} + \ln(1 - \frac{1}{\exp(\Theta_{E2}/T)}) \right], \quad (5)$$

where R is gas constant, m_B , m_{E1} and m_{E2} are degrees of freedom and $m_B+m_{E1}+m_{E2} \ge 3n$ and n is number of atoms, \Box_B , \Box_{E1} and \Box_{E2} are the characteristic temperatures Bose and Einstein parts, $b=1/[\exp(g)-1]$, $g=d\ln[1+\Box_B/(Td)]$, d is fitting parameters for heat capacity near 0 K. If d=1 we have linear dependence of heat capacity, for d=4 this dependence approaches Debye capacity and we have Einstein capacity.

From (5) we can obtained the internal energy and the heat capacity at constant volume:

$$E_{th} = m_B R\left(\frac{\Theta_B}{2} + \frac{\Theta_B T db}{T d + \Theta_B}\right) + m_{E1} R\left[\frac{\Theta_{E1}}{2} + \frac{1}{\exp(\Theta_{E2}/T) - 1}\right]$$

$$+ m_{E2} R \left[\frac{\Theta_{E2}}{2} + \frac{1}{\exp(\Theta_{E2}/T) - 1} \right],$$
 (6)

$$C_{V th} = m_B R \left(\frac{\Theta_B d}{T d + \Theta_B} \right)^2 b (1/d + 1 + b) +$$

$$+ m_{E1} R (\Theta_{E1}/T)^2 \frac{\exp(\Theta_{E1}/T)}{\left[\exp(\Theta_{E1}/T) - 1\right]^2}$$

$$[\exp(\Theta_{E1}/T) - 1]^{2} + m_{E2} R(\Theta_{E2}/T)^{2} \frac{\exp(\Theta_{E2}/T)}{[\exp(\Theta_{E2}/T) - 1]^{2}}.$$
 (7)

From (5) we have the thermal pressure $P_{th} = -(\partial F_{th}/\partial V)_T$:

$$P_{th} = \tilde{a}_R E_R / V + \tilde{a}_{E2} E_{E1} / V + \tilde{a}_{E2} E_{E2} / V$$
, (8)

where E_B , E_{E1} and E_{E2} are Bose and Einstein parts of the inner energy γ_B , γ_{E1} and γ_{E2} are the Gruneisen parameters $\gamma = -(\partial \ln \Box/\partial \ln V)_T$, and from (8) can be found $K_{Tth} = -V(\partial P_{th}/\partial V)_T$:

$$K_{Tth} = P_{thB} (1 + \gamma_B + q_B) - \gamma_B^2 TC_{VB} / V + + P_{thE1} (1 + \gamma_{E1B} + q_{E1}) - \gamma_{E1}^2 TC_{VE1} / V + P_{thE2} (1 + \tilde{a}_{E2} + q_{E2}) - \gamma_{E2}^2 TC_{VE2} / V ,$$
 (9)

where P_{thB} , P_{thE1} and P_{thE1} , C_{VB} , C_{VE2} and C_{VE2} are Bose and Einstein parts of thermal pressure and heat capacity at constant volume, q_B , q_{E1} in q_{E2} are logarithmic derivatives $q = (\partial \ln y / \partial \ln V)_T$.

At high temperatures the anharmonicity should be taken into account. According [1], the anharmonic part of the free energy is proportional to the square of temperature:

$$F_a(V) = -a(V)T^2, \tag{10}$$

where a(V) is the anharmonic parameter. From this we have:

$$E_a = aT^2$$
, $S_a = 2aT$ и $C_{Va} = 2aT$, $P_a = \gamma_a E_a/V$, $\gamma_a = (\partial \ln a/\partial \ln V)_T$, $K_{Ta} = P_a(1-\gamma_a)$. (11)

We assume that $\gamma_a = \text{const.}$

From the slop at constant volume

$$(\partial P/\partial T)_{V} = \gamma_{B} C_{VB} / V + \gamma_{E2} C_{VE1} / V + + \gamma_{E2} C_{VE2} / V + \gamma_{a} C_{Va} / V$$
(12)

we can found thermal expansion coefficient $\alpha = (\partial P/\partial T)_V/K_T$, where $K_T = K_T(V) + K_{Tth} + K_{Ta}$.

The volume dependence of the Gruneisen parameters and characteristic temperatures written in classical form:

$$\gamma = \gamma_0 x^q$$
 и $\Theta = \Theta_0 \exp[\gamma_0 (1-x^q)/q]$. (13) Pressure, enthalpy and the Gibbs energy may be found from $P = P(V) + P_{th} + P_a$, $H = E + PV$, $G = F + PV$.

We assume that thermodynamic values of the Gruneisen parameter is true for any conditions

$$\gamma = \alpha K_T V/C_V$$
 и $\gamma = \alpha K_S V/C_P$, (14) where K_S is adiabatic bulk modulus and C_P is heat capacity at constant pressure.

The offered model allows to calculate any thermodynamic functions at given temperature and volume or pressure if we know all fitting parameters: $U_0, V_0, K_0, K', d, m_B, m_{E1}, m_{E2}, \Theta_{B0}, \Theta_{E10}, \Theta_{E20}, a_0,$ γ_{Bo} , γ_{E1o} , γ_{E2o} , γ_a , q_B , q_{E1} , q_{E2} . The fitting parameters may be obtained by optimization, that is, by solving the set of equations using the weighted least squares method and experimental data on heat capacity at constant pressure, relative enthalpy, volume, thermal expansion coefficient, bulk moduli and other data. The offered set of equations can be simplified by suspecting, that the Gruneisen parameter and parameter q is identical for all thermal parts, that essentially reduce number of fitting parameters. Now we have a following set of fitting of parameters: U_0 , V_0 , K_0 , K', d, m_B , m_{E1} , m_{E2} , Θ_{B0} , Θ_{E10} , Θ_{E20} , a, γ , γ_a , q.

- Zharkov V.N., Kalinin V.F. Equation of state for solids at high pressures and temperatures // New York: Consultants Bureau. 1971. 257 p.
- Zharkov V.N. On the dependence of the coefficient of thermal expansion on density // Phys. Earth Planet. Inter. 1998. Vol. 109. No. 1–2. P. 79–89.
- 3. Pankov, V.L., Kalinin, V. A., and Kalachnikov, A. A. Phase Relations of the MgO–SiO₂ and FeO–SiO₂ systems and their implication for the properties and composition of the mantle. Izvestiya, Physics of the Solid Earth, 1996, 32, 484–495.
- 4. *Polyakov, V.B. and Kuskov, O.L.* Selfconsistent model for the calculation of thermoelastic and caloric properties of minerals. Geochimia, 1994, 1096–1122 (in Russian).
- Anderson, O.L. Equations of state of solids for geophysics and ceramic science. New York: Oxford University Press. 1995. 405 p.
- 6. *Hama J., Suito K.* Thermoelastic properties of periclase and magnesiowustite under high pressure and high temperature // Phys. Earth Planet. Interiors. 1999. V. 114. P. 163–179.
- Dubrovinskaya N.A., Dubrovinsky L.S., Saxena S.K. Systematics of thermodynamic data on solids: Thermochemical and pressure-volumetemperature properties of some minerals // Geoch. Cosmoch. Acta. 1997. V. 61. No. 19. P.4151-4158.
- 8. *Gerya, T.V., Podlesskii, K.K., Perchuk, L.L., Swamy, V., and Kosyakova, N.A.* Equation of state of minerals for thermodynamic databases used in petrology. Petrology, 1998, 6, 511–526.
- 9. *Dorogokupets P.I.* Thermodynamic functions at zero pressure and their relation to equations of state of minerals // Amer. Mineral. 2000. Vol. 85. No. 2. P. 329-337.
- Dorogokupets, P.I., Ponomarev, E.M., and Melechova, E.A. Optimization of experimental data on the heat capacity, volume, and bulk moduli of minerals. Petrology, 1999. 7. 574–591.
- 11. Anderson O.L., Zou K. Formulation of the thermodynamic functions for mantle minerals: MgO as an example // Phys. Chem. Minerals. 1989. Vol. 16. P. 642–648.
- 12. Kut'in A.M., Pyadushkin D.V., Russian Journal of Physical Chemistry, 1998, Vol. 72, No. 10, 1567-1572.
- Thompsen L., Anderson O.L. On the hightemperature equation of state of solids // J. Geophys. Res. 1969. Vol. 74. No. 4. P. 981–991.
- 14. *Poirier J.-P., Tarantola A.* A logarithmic equation of state // Phys. Earth Planet. Inter. 1998. Vol. 109. No. 1–2. P. 1–8.